## Computational and experimental studies of iron-bearing carbonates and silicate glasses at lower mantle pressures

Natalia Solomatova<sup>\*†1</sup>, Jennifer Jackson<sup>1</sup>, Paul Asimow<sup>1</sup>, Wolfgang Sturhahn<sup>1</sup>, George Rossman<sup>1</sup>, and Mathieu Roskosz<sup>2</sup>

<sup>1</sup>Division of Geological Planetary Sciences, California Institute of Technology, Pasadena, CA, United States – United States

<sup>2</sup>IMPMC-Muséum National d'Histoire Naturelle-CNRS, Paris, France – Muséum National d'Histoire Naturelle (MNHN) – France

## Abstract

Decomposition of carbonates may be responsible for creating silicate melts within Earth's mantle by lowering the melting temperature of surrounding rock. Characterizing the stability of carbonate phases is therefore a necessary step towards understanding the transport and storage of carbon in Earth's interior. Dolomite is one of the major mineral forms in which carbon is subducted into Earth's mantle. Although iron-free dolomite is expected to break down upon compression, high-pressure polymorphs of iron-bearing dolomite may resist decomposition. Using a genetic algorithm that predicts crystal structures, we found a monoclinic phase with space group C2/c that has a lower energy than all previously reported dolomite structures at pressures above 15 GPa, where the substitution of iron for magnesium stabilizes monoclinic dolomite with respect to decomposition at certain pressures of the lower mantle. Thus, an iron-bearing dolomite polymorph may be an important carbon carrier in regions of Earth's lower mantle.

The depth at which carbonates will decompose is dependent on the age, temperature and density of subducting slabs. Decarbonation reactions may lower the melting temperature of surrounding rocks to produce silicate melts. In regions of the mantle where silicate melts may exist, it is important to understand the physical properties and dynamic behavior of the melts because they affect the chemical and thermal evolution of its interior. Composition, degree of polymerization, and iron's spin state affect such properties. The behavior of iron in silicate melts is poorly understood but, in some cases, may be approximated by iron-bearing glasses. We measured the hyperfine parameters of iron-bearing rhyolitic and basaltic glasses up to  $_120$  GPa and  $_200$  GPa, respectively, in a neon pressure medium using synchrotron M'ossbauer spectroscopy. The spectra for rhyolitic and basaltic glasses are well explained by three high-spin Fe2+-like sites with distinct quadrupole splittings, reflecting the influence of evolving coordination environments with pressure. With the assumption that coordination environments in silicate glasses may serve as a good indicator for those in a melt, this study suggests that ferrous iron in chemically–complex silicate melts likely exists in a high-spin state throughout most of Earth's mantle.

\*Speaker

 $<sup>^{\</sup>dagger}$ Corresponding author: nsolomat@caltech.edu